

PETROLOGIC AND GEOPHYSICAL CHARACTERISTICS OF THE LITHOSPHERE AND THE LITHOSPHERE-ASTHENOSPHERE BOUNDARY THROUGH TIME

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ABSTRACT

The lithosphere-asthenosphere boundary (LAB) is an interface that reflects the fundamental evolution of the Earth into major geochemical reservoirs driven by global thermal processes. The Earth's lithospheric mantle is non-convecting (in contrast to the asthenosphere) and does not mix or homogenise easily: therefore it carries a geochemical imprint of major melting events and subsequent fluid fluxes (metasomatism). As a generalisation, lithospheric mantle has undergone significant (probably multiple) melting and is depleted in "basaltic" components such as Fe, Al, Ca and Ti. Asthenosphere by contrast, is "fertile" (rich in basaltic components) and geochemically more homogeneous than lithosphere because it is convecting.

The depth to the LAB varies significantly with the tectonothermal age of lithosphere domains and is a function of the geochemical nature of the lithospheric mantle column and its thermal state which determine its density and rheological response. Mantle xenoliths and their disaggregated minerals (especially garnet and chromite) brought to the surface in basalts and kimberlites, can be used to construct sections of the rock types and their spatial distribution with depth in the lithospheric mantle, to estimate empirical palaeogeotherms, and to locate the depth to the LAB (O'Reilly and Griffin, 1996). This geochemically determined depth coincides with estimations of depth using geo-

physical datasets, especially seismic refraction and magnetotelluric.

We have also established (Griffin et al., 1998; 1999) that the mean geochemical composition of the subcontinental lithospheric mantle beneath terranes with Archean, Proterozoic and Phanerozoic crustal tectonothermal ages shows an episodic secular and apparently irreversible change with time. Younger lithospheric mantle is progressively less depleted in "basaltic" components. Using average mineral compositions measured from xenoliths, these compositions have been recast into mineral assemblages (modes), and used to derive the mean density and seismic velocities for mantle lithosphere of different ages (Figure 1). Archean lithospheric mantle is much less dense than Phanerozoic mantle, but has higher seismic velocities at any temperature; Proterozoic mantle lies between these two end-members.

This correlation of mantle composition with crustal age has important implications:

- The crust and the underlying lithospheric mantle must have formed quasi-simultaneously, and they typically have stayed coupled for periods of aeons
- There must have been major changes in Earth's geodynamics about 2.5 Ga ago, which changed the way continents and their roots are formed

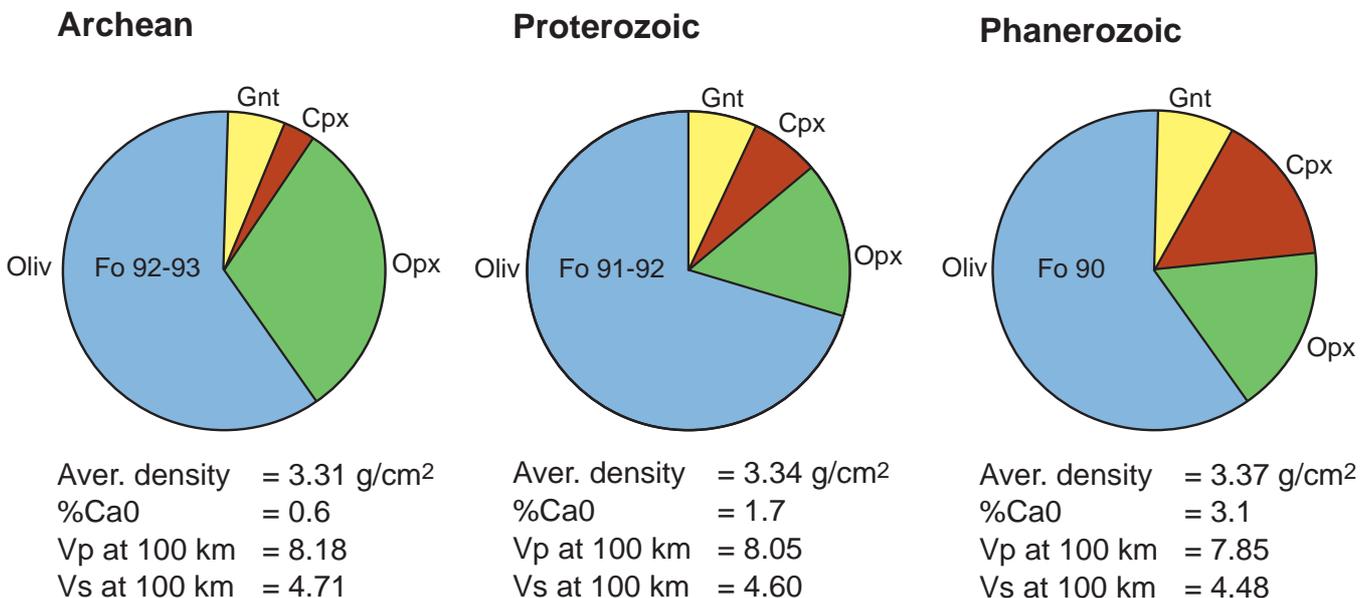


Fig. 1 - Differences in mantle mode, composition and physical properties for Archean, Proterozoic and Phanerozoic lithospheric mantle.

- Since 2.5 Ga, the continent-forming processes have evolved toward lower degrees of partial melting; this suggests a link to the secular cooling of Earth

Density and seismic measurements and calculations for these lithospheric mantle types reveal an increase in density and decrease in V_p and V_s with decreasing age. Geothermal gradients increase concomitantly and average depths to the LAB decrease from about 220 km to 180 km to 100 km respectively. Density and thermal differences in these distinctive mantle volumes determine the relative buoyancy of the specific lithospheric mantle column relative to asthenosphere. Archean mantle is depleted (Mg-rich) and buoyant, but seismic velocities are higher due to olivine's property of higher V_p and V_s with increasing Mg content (but decreasing density) enhanced by the lower geothermal gradient. Phanerozoic lithosphere is fertile and thus relatively dense despite the typically higher geothermal gradients. Phanerozoic SCLM with relaxed geothermal gradients is denser than the asthenosphere at the same depth, and hence is gravitationally unstable. Thus the geochemical composition and thermal state of lithospheric mantle may control the location of the LAB through time and result in the different geophysical signatures.

Archean lithosphere is refractory, and too buoyant to be "delaminated" by gravitational or thermal effects and this may explain the longevity and stability of Archean keels.

Replacement of Archean lithosphere must involve mechanical dispersal (eg during rifting) and may be accompanied by thermal and chemical erosion as hot, fertile asthenospheric material rises; the tectonic consequences at the surface include uplift, magmatism and basin formation during subsequent thermal relaxation. Such lithosphere replacement events may be important in producing fluids and structures seen in the lithospheric mantle as metasomatic episodes.

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